Low-temperature deep geothermal systems: an overview



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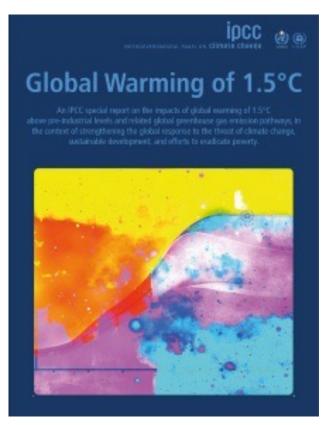


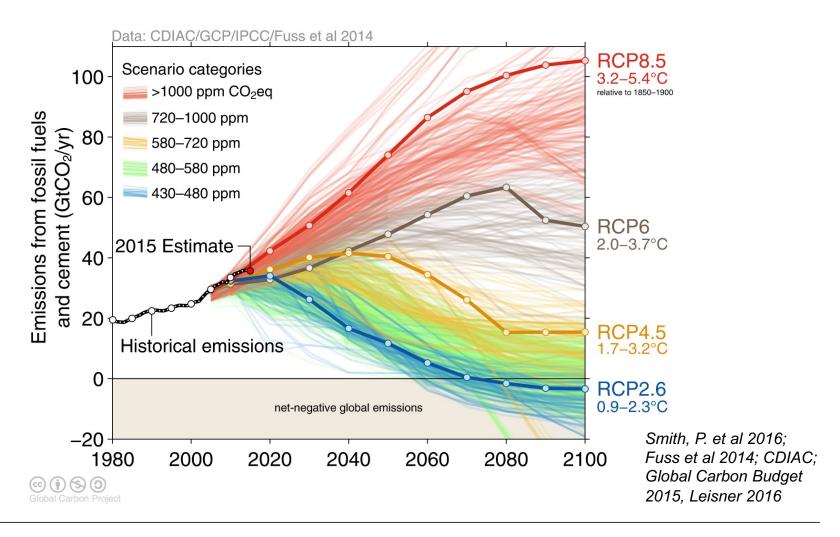
H2020 DEEPEGS project: IDDP-2 well Reykjanes peninsular 4665 m

Intergovernmental panel on climate change: Scenarios to reach COP21 goal



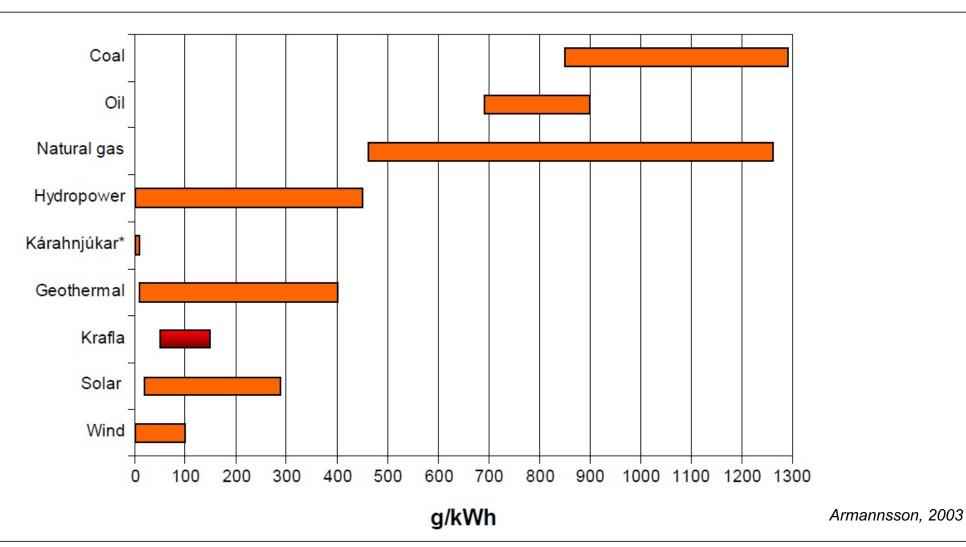
- IPCC Global Warming report
 - http://www.ipcc.ch/report/sr15/





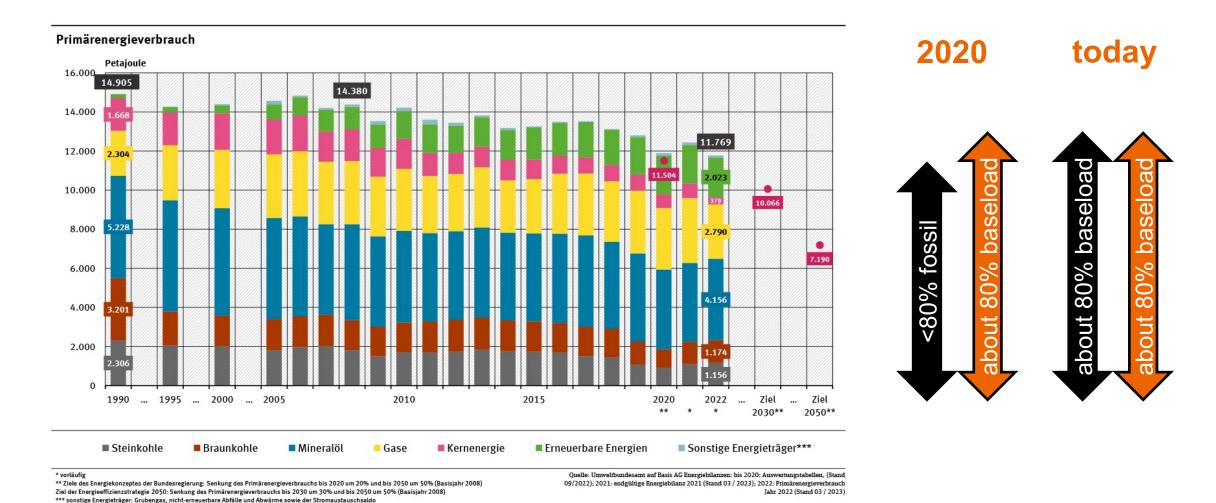
CO₂ emission from reservoir fluids





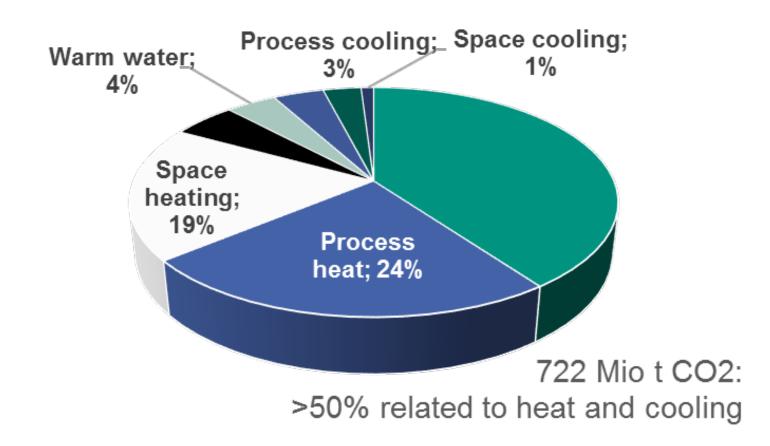
Share of fossil and baseload energy in the German mix





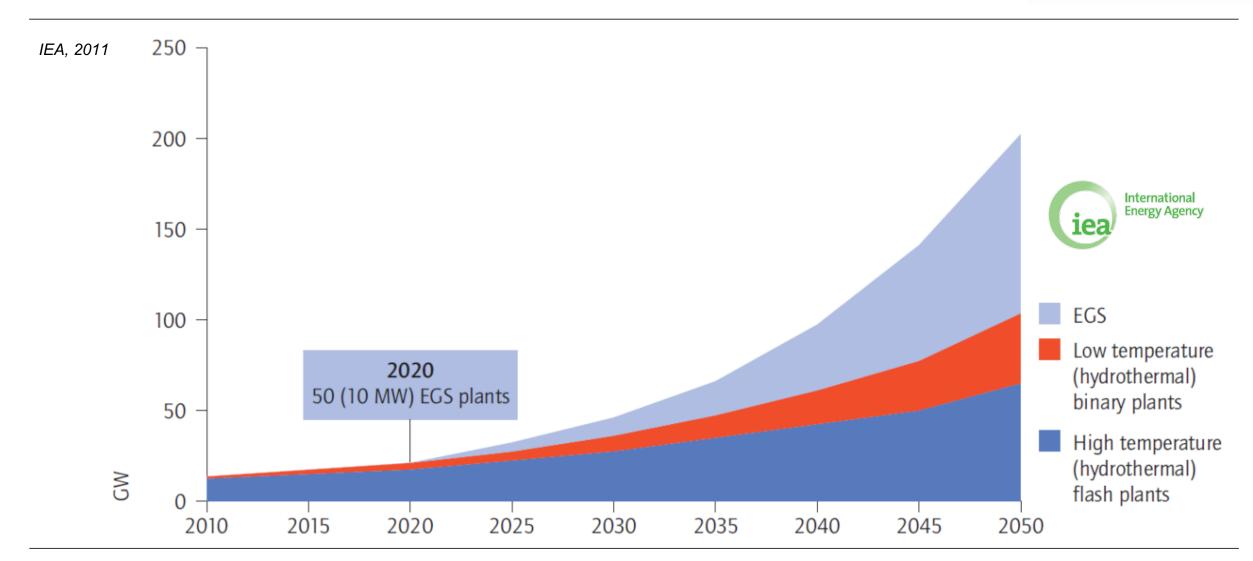
CO₂ emissions in the German energy sector





Global geothermal power potential





Geothermal energy utilisation in the history



■ 6200 BC Hasan Dag volcano VII level of the Catal Hüyük village in Anatolia (Turkey, Mellaart, 1967)

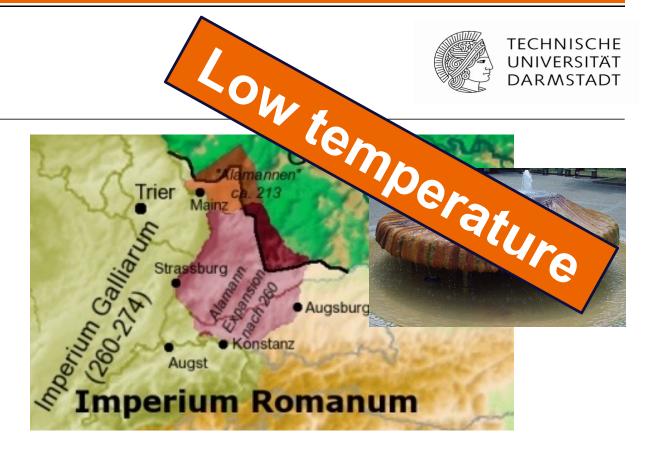
 6th century BC Holy fountain of Giara Santa Vittoria (Sardinia, Italy, www.laghienuraghi.it) sparkling volcanic water





Roman culture to medieval time

- Thermal spa Wiesbaden (Germany)
 - 77 AD Naturalis historia (Plinius the Older)
 - 121 AD Aquae Mattiacorum, Roman settlement
 - 259/260 AD Destruction by Teutons
 - 306-337 AD Recapture of the settlement by the Romans (only settlement of Romans Est of the Rhine river)
- Italy after the Roman Empire
 - **1400 AD** Balneotherapy and the use of geothermal products become even more important locally in Tuscany.
 - 1472 End of the "Guerra delle Allumiere" between Volterra and Florence over geothermal manifestations in the "Regione boracifera" (now Larderello). The right of use passes to the Medici family.



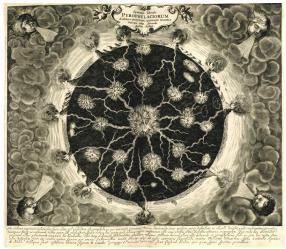


1664-1665 AD during the renaissance Geocosmus (Athanasius Kircher)





Hydrophylacia



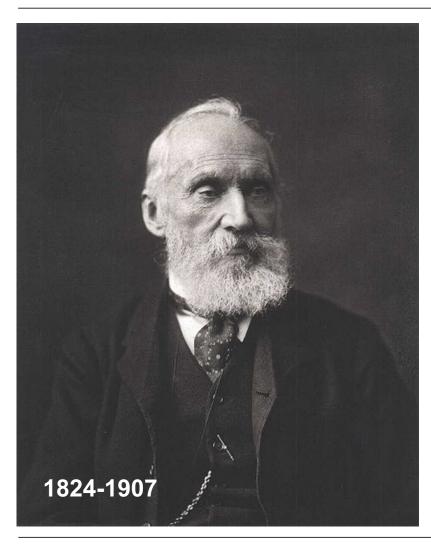
Phyrophylacia



Slide 9

1880 AD Lord Kelvin





- Age of the Earth
 - Several 100 millions of years
 - His assumption: the heat flux at the Earth's surface represents the remaining heat from the initial phase of the Earth

$$\frac{1}{A}\frac{dQ}{dt} = \frac{\dot{Q}}{A}$$

■ Fourier's Law of Heat Conduction (1822)

$$T = \frac{q_0}{k} \cdot z + T_0$$

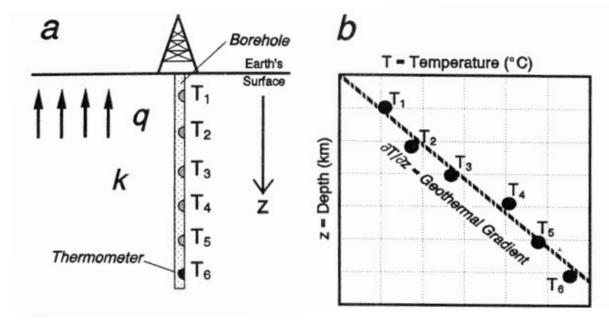
- Boundary condition
 - surface temperature T₀ and
 - heat flux at the surface q₀

Determination of heat flux density

TECHNISCHE UNIVERSITÄT DARMSTADT

 Heat flux density from depthdependent temperature measurements

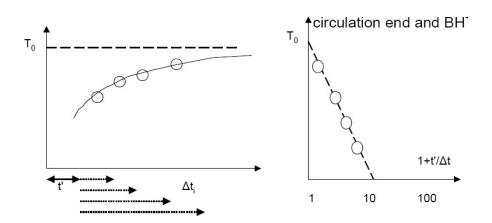
$$T = \frac{q_0}{k} \cdot z + T_0$$



Determination of temperatures

$$T(t) = T_0 - k \cdot \log \left(1 + \frac{1}{\Delta t} \right)$$
urbed rock temperature

- T0 undisturbed rock temperature
- k constant depending on the λ of mud and surrounding rock
- t' cooling period (duration of circulation at bottom hole)
- Δt recovery period (time between circulation end and BHT measurement)
 (Lachenbruch and Brewer, 1959)



1904 First conversion of geothermal to electric energy



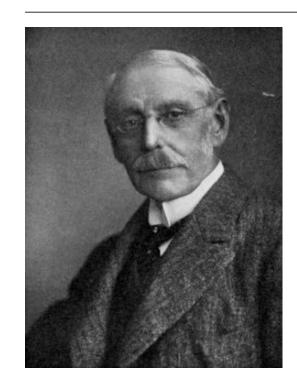
- 1904 First conversion of geothermal to electric energy by Conte Ginorio Conti at Larderello using a heat exchanger. The vapor operated a piston motor that connected to a dynamo of 10 kW lightening five lights.
- 1905 Installation of a piston motor connected to a 20 kW dynamo



1904 « Hellfire Exploration Project »



ਪion for



Sir Charles Parsons published his idea « Hellfire « Presidential Address » in the Annual Report of the better the Advancement of Science: Transaction of Section G, En 667-676):

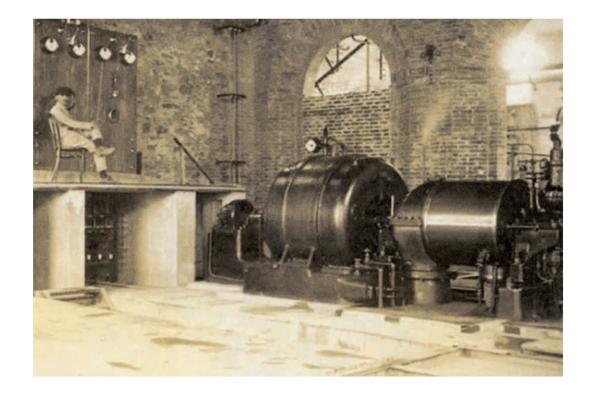
« The thermal system of the subsurface may be exploited by drilling to a depth of 20km. »

- He estimated
- A duration of construction of 85 years,
- a cost of 5 million pounds and a
- temperature of 600°C (assuming a geothermal gradient of 30°C per km)

Continues development in conventional systems



■ 1913 First geothermal plant with a installed capacity of 250kW in La Valle del Diavolo



■ **Today** about 700 MW installed capacity in 32 plants producing from three productive areas

The Geysers



■ The area is noted for numerous geysers and hot springs with surface temperature near 100°C. The area was originally settled by the **Pomos** and **Mayacmas Indians**.

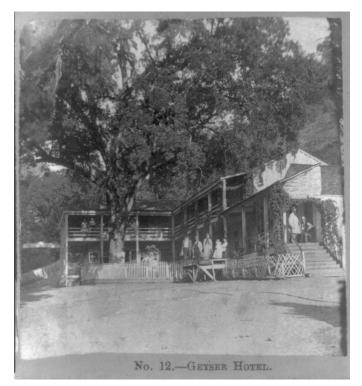
■ To the Native Americans this was "Tu-la-ha-lu-si, **the beautiful land**"; and the hot, spongy turf was

"Coo-lay-no-maock, the oven place", according to a local historian.

■ 1862 Samuel Brannan established a resort and spa.

Lund, 1995

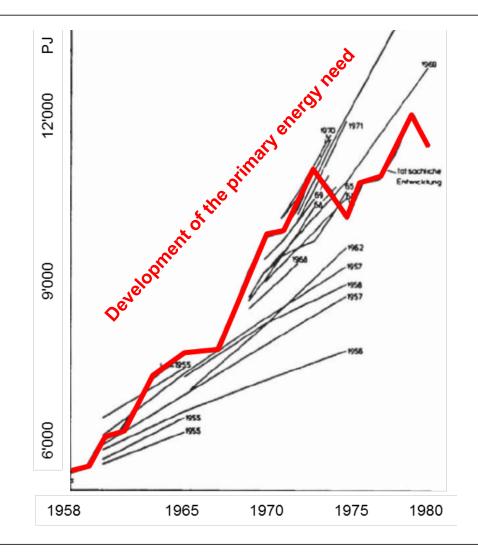
■ 1960 Pacific Gas and Electric began operation of a 11MW power plant.



Impact of oil crises on geothermal development

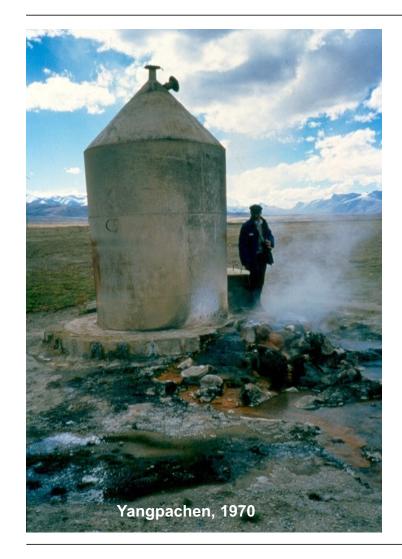


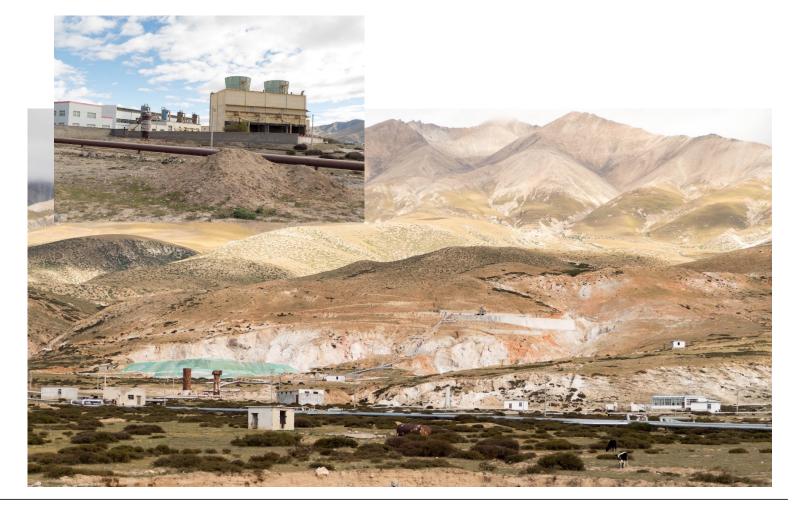
- Energy need between 1958-1980
 - Significant increase in the late 1960-ties
 - 1973 oil crisis: price increased 400%
 - 1979 oil crisis: price increased 100%



Geothermal energy in Tibet







Eva Schill / 19.11.24 - 17 -

Kola Super-Deep well KSDB-3



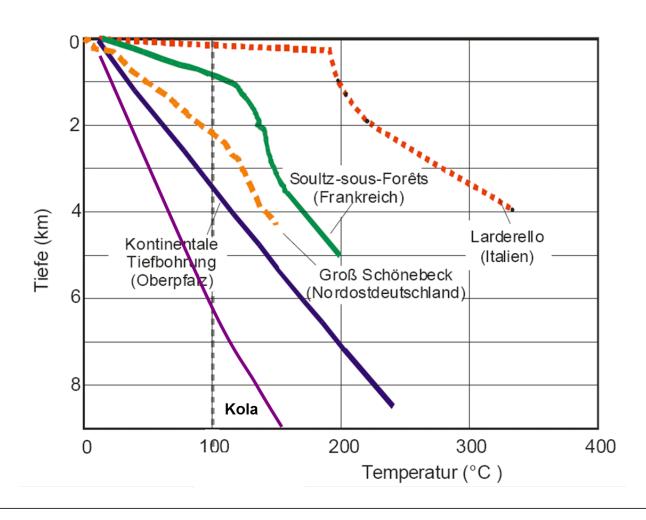
- 1970 Start of drilling of KSDB-3 in the northern part of the Pechengsk basin a thick volcanosedimentary series of mid-Proterozoic age and a gneissic basement
 - Target depth: 15 km
- **1975** T>120°C at a depth of 7'263m
- 1980 T>180°C in open-hole to a depth of 10'500m bis T>180°C
- 1989 Final depth of 12'261 m



Band width of geothermal wells in 2010



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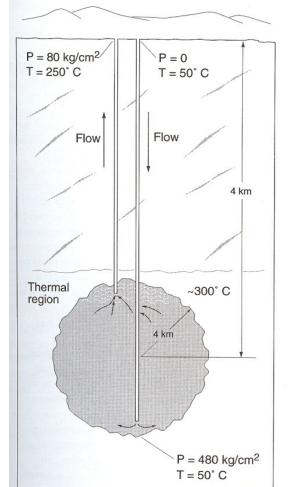


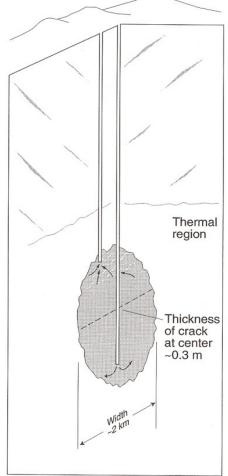
1970 Hot Dry Rock concept



- Hot Dry Rock programme USA (Los Alamos, Robinson 1971)
 - Artificial conductivity by stimulation

Rosemanowes Fenton Hill Soultz

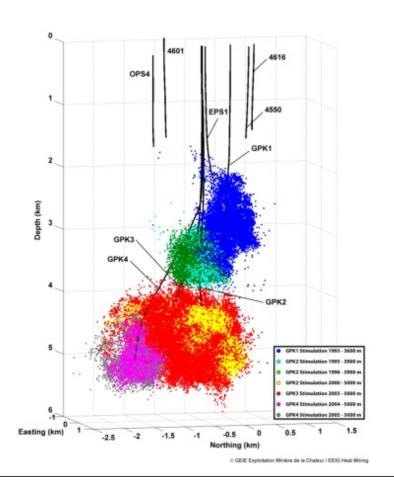




1987 Concepts for non-volcanic, low-temperature zones

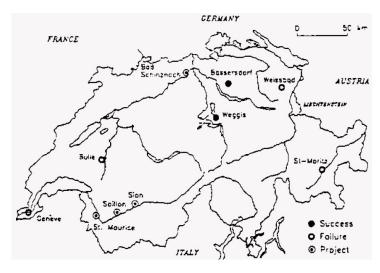


HDR to EGS



Low-temperature hydrothermal wells

- Federal programme « Risk guaranty » 1987-1998
- Project « Geothermoval » 1987-1995



Heat transport in low-temperature systems



Heat transport equation (without adiabatic effect)

$$\rho c_p \frac{\partial T}{\partial t} = \nabla (k \nabla T) + A - [\rho c_p]_f v_D \nabla T$$

Energy content of sources from the geosphere



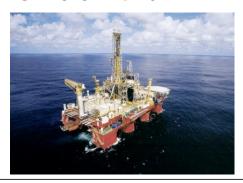
Heat value

■ Lignite: 1.7×10⁷ J kg⁻¹

■ Black coal: 3.0×10⁷ J kg⁻¹



■ Oil: 3.8×10⁷ J L⁻¹



Heat capacity

Thermal water:
 4.2×10³ J kg-1 K⁻¹

Rock:
 2.0×10³ J kg-1 K⁻¹



Cooling 1 kg of water or rock by 100K, provides 0.5-1% of the heat that is produced by burning the same amount of fossil oil or coal

$$c_p = \frac{\Delta E}{m \cdot \Delta T}$$

$$\Rightarrow \Delta E = c_p \cdot m \cdot \Delta T$$

$$\Rightarrow \frac{\Delta E}{t} = \Delta P = c_p \cdot \frac{\rho \cdot V}{t} \cdot \Delta T$$

$$\Rightarrow \Delta P = \rho c_p \cdot Q \cdot (T_{prod} - T_{inj})$$

Heat demand in Germany (2019)



Communities

Space heating

Water

Industrial demand

Process heat

Space heating

788 TWh/a

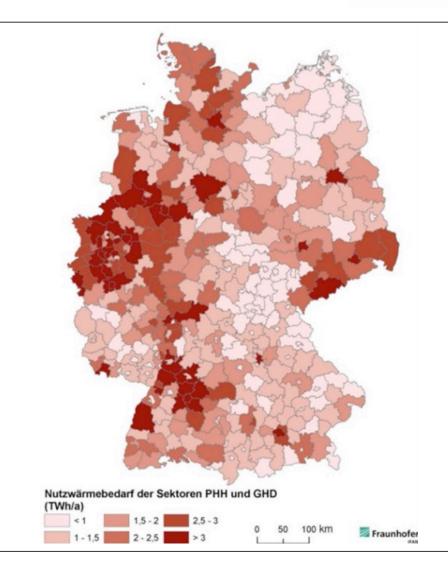
658 TWh/a (>2100 h/a)

130 TWh/a (8600 h/a)

604 TWh/a

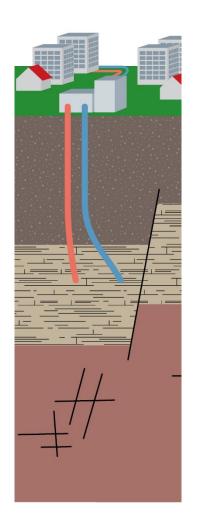
541TWh/a 8600 h/a)

63TWh/a (8600 h/a)



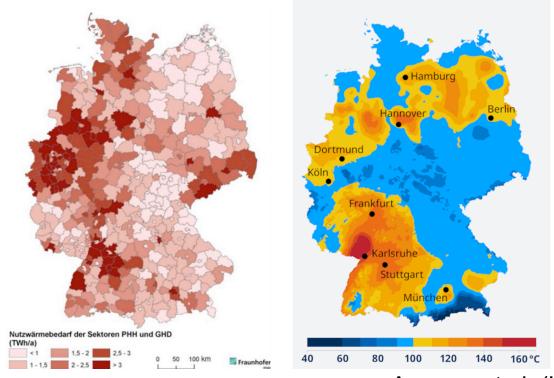
Hydrothermal Potential in Germany







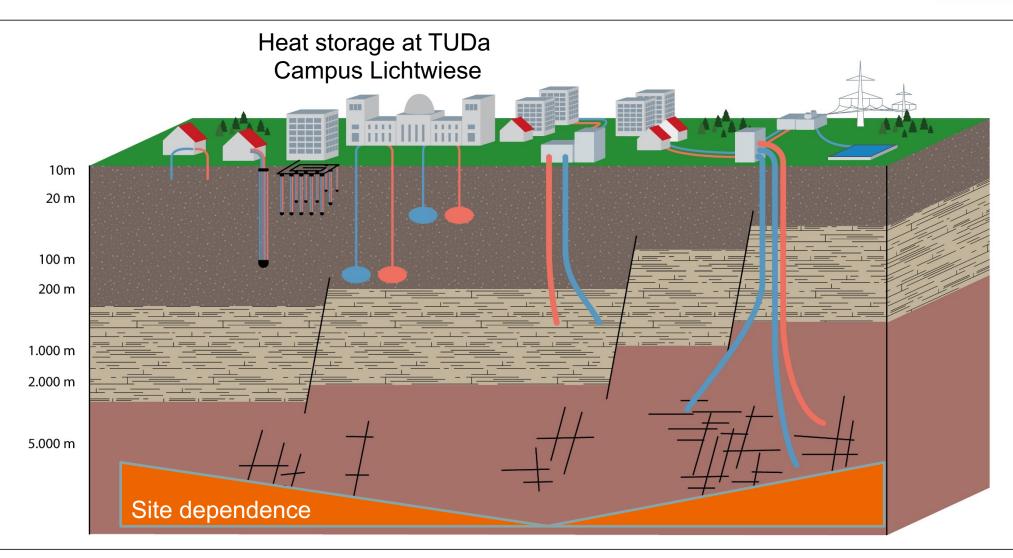
- Goal >2045 for hydrothermal ressources in germany
 - >300 TWh/a, i.e. 25% of the national heat demand



Agemar et al. (LIAG)

Geothermal technologies in Germany

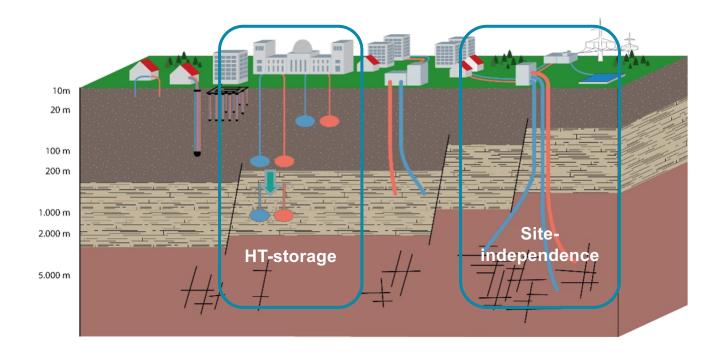


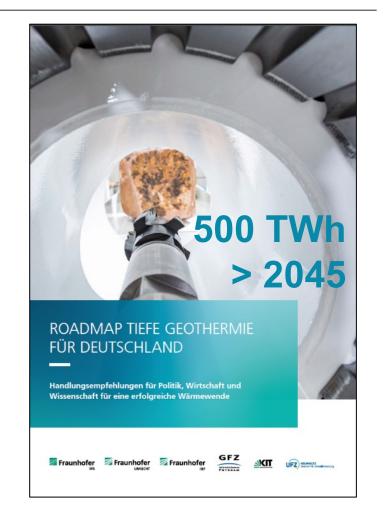


Geothermal potential including new developments



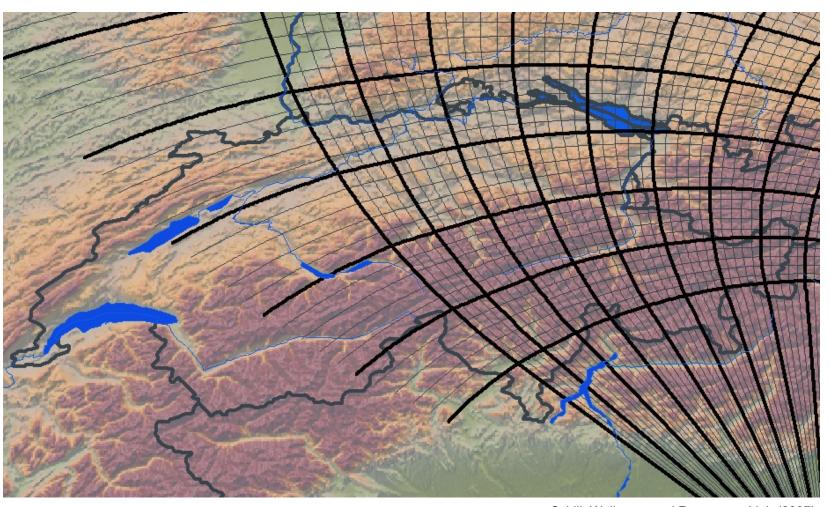
- Large-volume and high-temperature heat storage
- EGS Enhanced Geothermal Systems technology





Hydrothermal projects in and near the Alps





Schill, Wellmann and Regenauer-Lieb (2007)

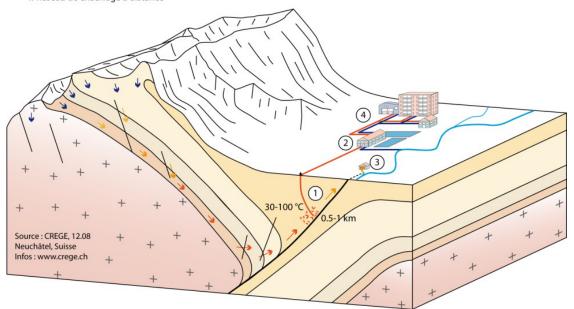
Alpine systems



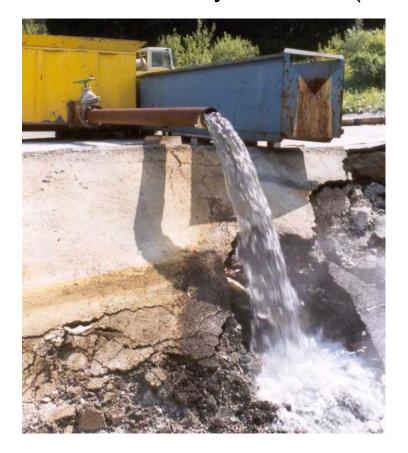
Typical concept of thermal utilisation

Chaleur des tunnels et hydrothermalisme

- 1. Forage de production
- 2. Centre thermal
- 3. Source thermale
- 4. Réseau de chauffage à distance



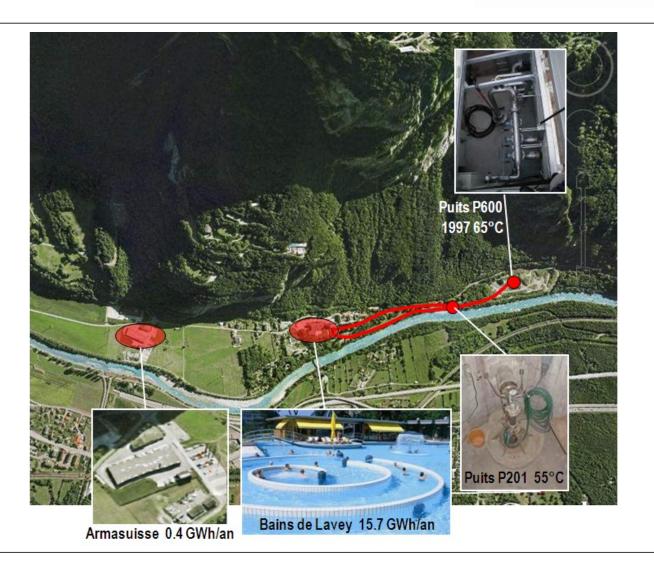
Production test Lavey-les-Bains (P600)



Spa / direct heat utilisation



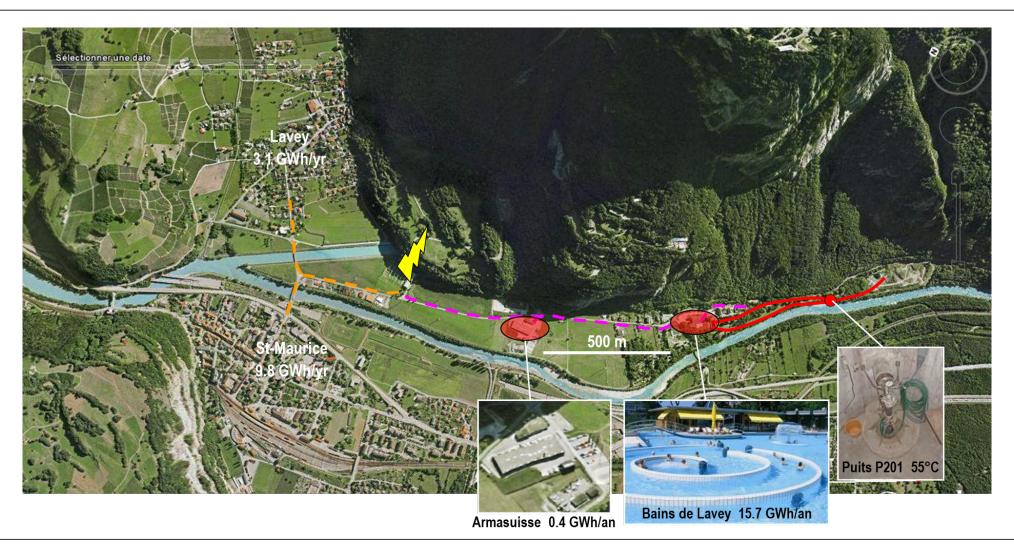
- P201 (201 m, 1973)
- P600 (600 m, 1997)
- Total flow rate 27 l/s
- Maximum temperature of 67°C



Planned extension failed



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Simplified Darcy's experiment



Steady state production rate

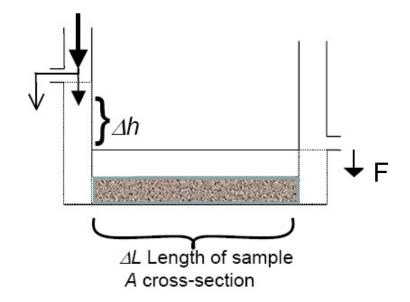
$$F = \frac{K \cdot A \cdot (h_1 - h_2)}{L}$$

 Volumetric flow rate: steady state flow rate per cross-sectional area

$$q = \frac{F}{A} = \frac{k \cdot \rho \cdot g}{\mu} \cdot \frac{dh}{dL} = K \cdot \frac{dh}{dL} = \nu_D \left[\frac{m}{S} \right]$$



$$\nu = \frac{\nu_D}{\frac{V_{connected\ void\ space}}{V_{total}}} = \frac{\nu_D}{\eta_{e\ (effective\ porosity)}}$$



 v_D = Darcy Fluid velocity [m s⁻¹]

 μ = Dynamic viscosity (fluid) [Pa s]

 ρ = Density (fluid) [kg m⁻³]

 $g = Gravity [m s^{-2}]$

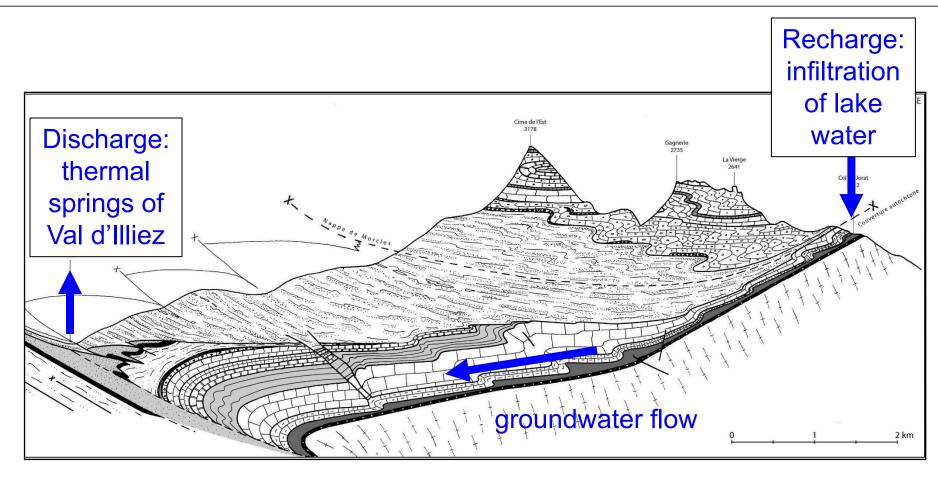
 $k = \text{Permeability } [\text{m}^2] \text{ or } [1 \text{ Darcy } D = 10^{-12} \text{m}^2]$

 $K = \text{Hydraulic conductivity } [\text{m s}^{-1}]$

F = Steady state production rate [m³ s⁻¹]

Hydrogeological interpretation of the geological profile Val d'Illiez





Pantet (2004)

Heat transport in low-temperature systems

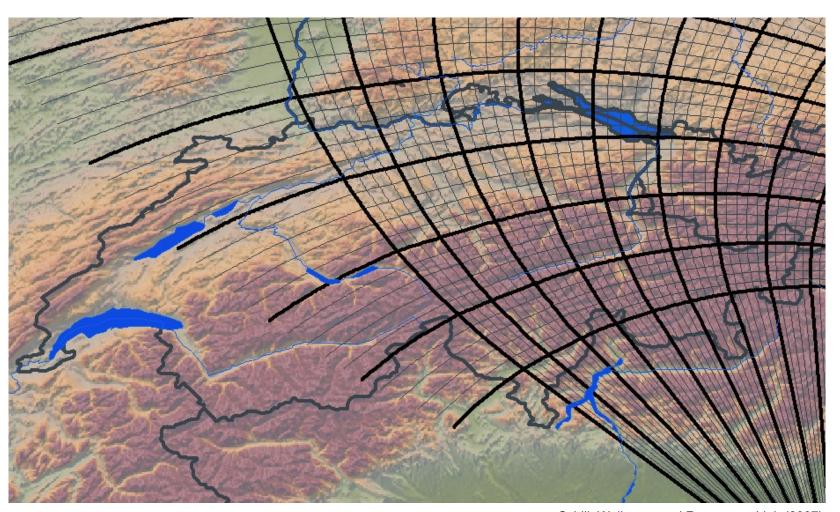


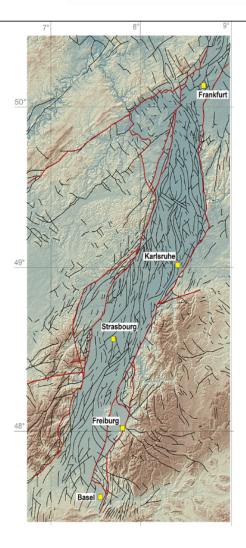
Heat transport equation (without adiabatic effect)

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Hydrothermal projects in and near the Alps





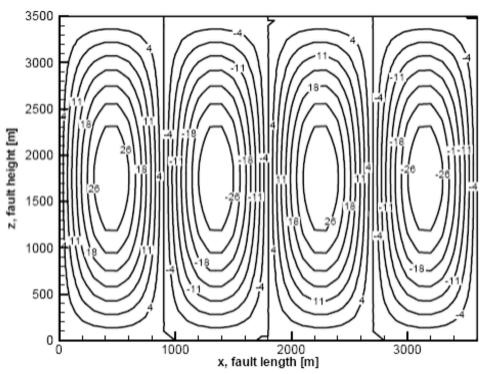


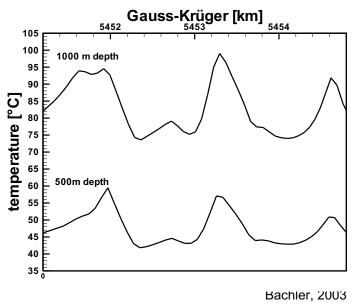
Schill, Wellmann and Regenauer-Lieb (2007)

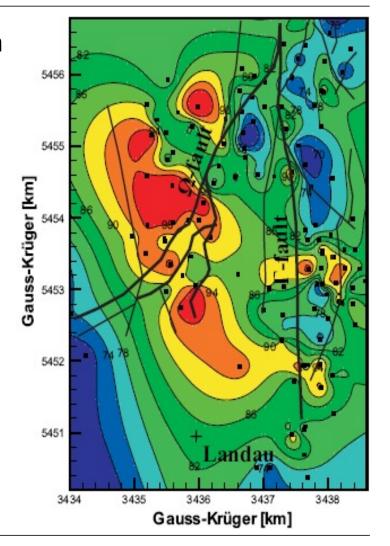
Indication for free (gravity-driven) convection in the URG



- Temperature distribution along fault zones at Landau in 1000m depth
 - Systematic temperature differences along Γ-fault
 - Analytical calculated temperature disturbance in the Γ-fault of about 15-20°C between lows and highs







(Critical) Rayleigh number



Ratio between hydraulic permeability and heat conductivity

$$Ra = \frac{\mathbf{k} \cdot \alpha \cdot \rho^2 c_p \cdot g \cdot \Delta T \cdot \Delta z}{\mu \cdot \lambda}$$

- Critical Rayleigh numbers
 - Free surfaces: Ra = $\frac{27}{4}$ π⁴ ≈ 657.51 (Lord Kelvin, 1916)
 - Rigid lower boundary: Ra = 1100.65
 - Rigid lower and upper boundaries: Ra = 39.48

k = hydraulic permeability

 α = thermal expansion

 ρ = density

 c_{p} = heat capacity of the fluid

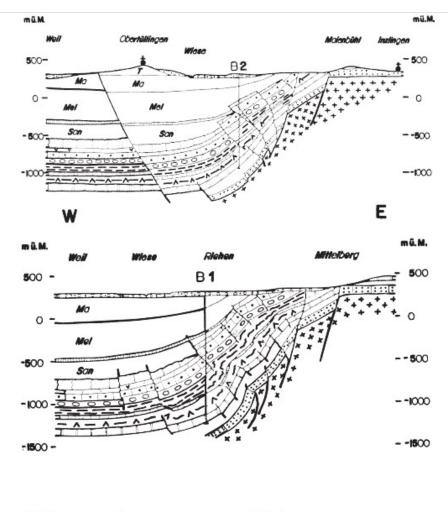
g = gravitational acceleration

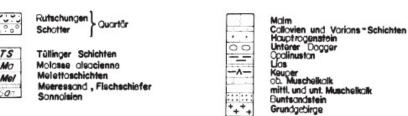
 μ = fluid viscosity

 λ = heat conductivity

Deep geothermal project at Riehen, Switzerland

- The sediments in the vicinity of the URG boundary fault
- highly factured
- small tectonic units [Gürler et al., 1987].
- The productive formation is the Upper Muschelkalk.
- In order to obtain maximum production, the highly fractured zones near the boundary fault were selected for the doublet system [Boissavy & Hauber, 1994].



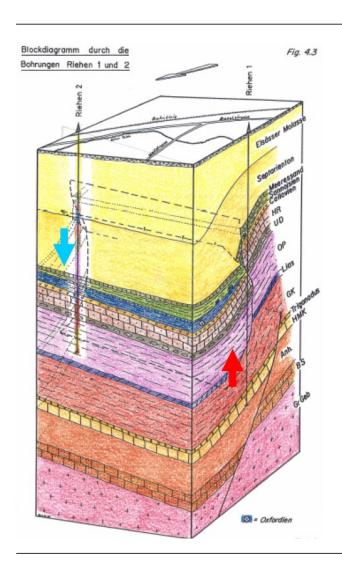




[Baudepartement Basel-Stadt 1989]

Naturally fractured reservoir Riehen (Switzerland)



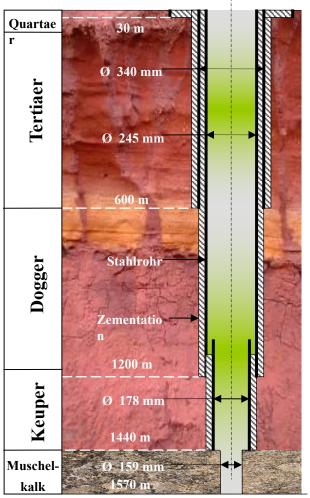


- 1988 Well RB1 (producer) and RB2 (injector)
- 1997 First trans-border heating system between Germany and Switzerland
- 2006 Riehen-Plus: Enhancement of the production rate from 18 to 23 L/s
- Production from the Upper Muschelkalk using a doublet system. Strong mineralisation of the brine impedes a singlet system. The well Riehen1 is deeper and thus warmer as Riehen 2.
- The reinjection guarantees the maintainance of the pressure (sustainable reservoir management)

Production well RB1 (Bachtelenweg)





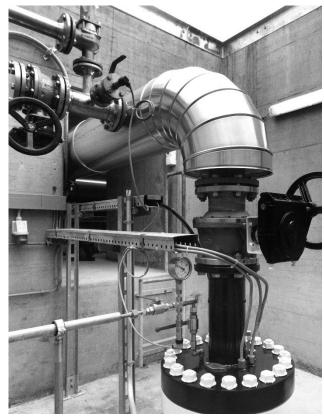




Wellhead under construction and in use









Deep geothermal project at Riehen

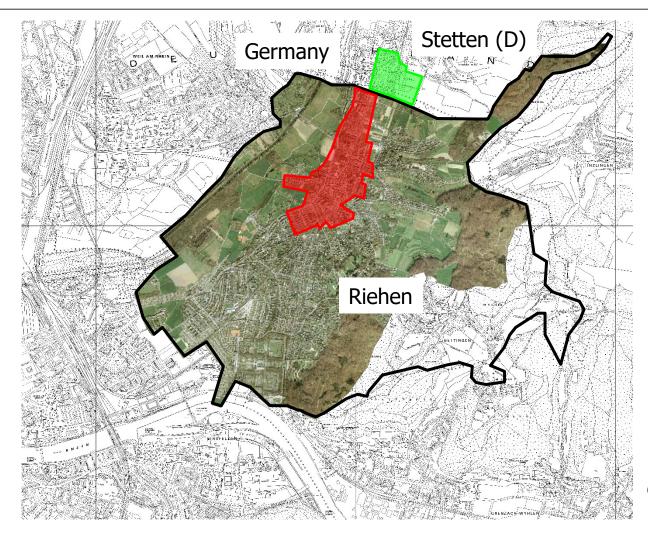


	Riehen 1	Riehen 2
Depth of Upper Muschelkalkes:		
Trigonodusdolomit	1444 - 1464 m	1123 - 1158 m
Hauptmuschelkalk	1464 - 1536 m	1158 - 1223 m
	2 m	9 m
	(1 fracture)	(3-4 fractures)
Pressure:		
	300m a.s.l.	300m a.s.l.
	+ 25 m artesic	+ 15 m artesic
Temperature in the Aquifer:		
	66,4 °C	52,2 °C
Distance between wells:		
	960 m	

Direct use



- Service aera as per 2009
 - Production rate 18'300 kW
 - Customers 202
 - Production temperature 65°C
 - Injection temperature 29°C
 - Production 18 l/s (~65 m³/hr)
 - Sales of heat 26'000 MWh/a
 - Heating district 21.4 km
 - The sales of geothermal heat corresponds to the use of 3'000'000 liters of oil





Concept of district heating at Riehen



